



## *Enoploclytia tepeyacensis* n. sp. (Crustacea, Decapoda, Erymidae) from the Cretaceous (Campanian) of Coahuila, NE Mexico

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### Abstract

The erymid lobster *Enoploclytia tepeyacensis* n. sp., is reported for the first time in the early Campanian beds of the Austin Group (Dessau Chalk) near Jiménez, Coahuila, NE Mexico. The new species represents the most complete record of *Enoploclytia* in North and Central America. The third report of *Enoploclytia* in Mexico is documented.

Keywords: Crustacea, Decapoda, Erymidae, Campanian, Cretaceous, Mexico.

### Resumen

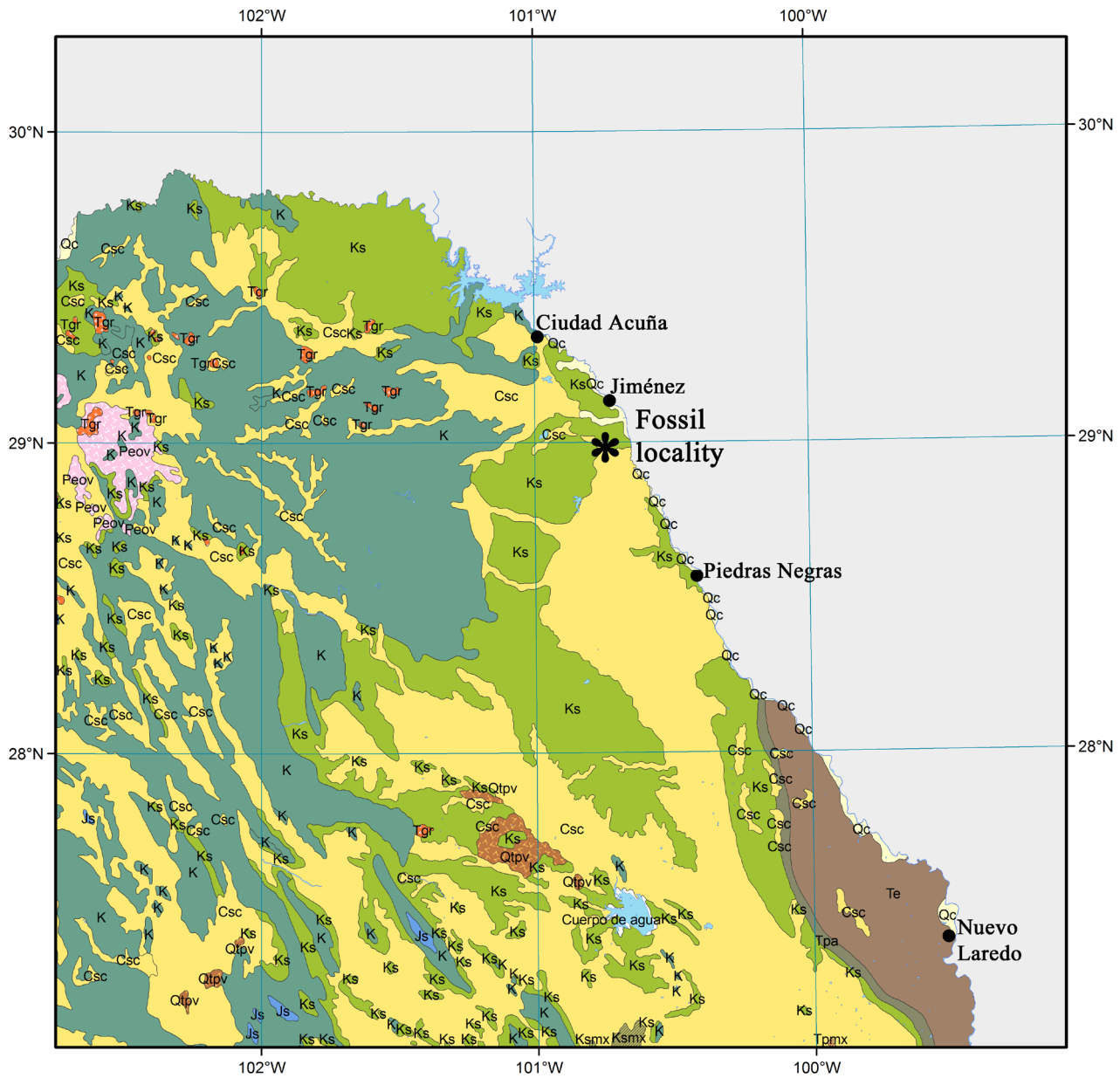
La langosta erímida *Enoploclytia tepeyacensis* n. sp. se reporta por vez primera en las capas del Campaniano inferior del Grupo Austin (Dessau Chalk) en las inmediaciones de Jiménez, Coahuila, NE de México. La nueva especie representa el registro más completo de *Enoploclytia* en América del Norte y América Central. Se documenta el tercer reporte de especies del género *Enoploclytia* en México.

Palabras Clave: Crustacea, Decapoda, Erymidae, Campaniano, Cretácico, México.

### 1. Introduction and geological setting

Rocks of the Dessau Formation (Campanian, Austin Group) crop out near Jiménez, Coahuila, NE Mexico (Figure 1). The Dessau Formation has an approximate thickness of 30 m in the study area, where invertebrate fossils including ammonoids and echinoids are found in association with the arthropod described here. Several reports deal with the stratigraphy of the Dessau Formation (Stephenson, 1936; Young and Marks, 1952; Young, 1963; Young and Woodruff, 1985; Lundquist, 2000; Cobban *et al.*, 2008; among others). Housh (2007) wrote, “The Dessau Chalk is a white chalk characterized by a *Phrygia aucella*

(Roemer, 1849) biostrome. The lower portion, up to the *Phrygia aucella* layer comprises the “*Gryphaea*” *aucella* zone of Young and Marks (1952); the overlying *Exogyra leviuscula* zone comprises the upper portion. The Dessau Chalk has a thickness of approximately 29 m in the Austin area. The Dessau Chalk is disconformably overlain by the Burditt Marl Formation. Carrasco (1969) reported the large ammonoid *Parapuzosia bosei* Scott and Moore, 1928, from a locality nearby where the specimens here reported were found, along with *Menavites (Menavites) internodosus* (Renz, 1936) and *M. (M.) desinodosus* (Renz, 1936), and suggested an early Campanian age for the Dessau beds on this region.



- Qc Quaternary alluvium and soils
- Qtpv Pliocene to Quaternary continental volcanic igneous rocks
- Csc Miocene to Quaternary fluviolacustrine deposits
- Peov Eocene to Oligocene continental volcanic igneous rocks
- Te Eocene marine sedimentary rocks
- Tpa Paleocene marine sedimentary rocks
- Tgr Tertiary intrusive igneous rocks (granite and gabroides)
- Ks Upper Cretaceous marine sedimentary rocks
- Ksmx Upper Cretaceous mixed sedimentary rocks
- K Cretaceous marine sedimentary rocks
- Js Upper Jurassic marine sedimentary rocks

Figure 1. Geologic map of the fossil locality near Ejido El Tepeyac, Jiménez County, Coahuila. Modified from Servicio Geológico Mexicano.

## 2. Material

Two specimens, three-dimensionally preserved, assigned to *Enoploclytia tepeyacensis* n. sp. (Erymidae Van Straelen, 1925), are deposited at the Museo de Múzquiz, Coahuila (MUZ).

The systematic arrangement used in this paper follows the recent classification proposed by Schweitzer *et al.* (2010), using the following abbreviations: hcxp: height of carapace; lcxp: length of carapace; wcxp: width of carapace.

## 3. Systematic palaeontology

Superfamily Erymoidea Van Straelen, 1925

Family Erymidae Van Straelen, 1925

Genus *Enoploclytia* McCoy, 1849

Type species *Astacus leachii* Mantell, 1822, by monotypy

Fossil species see Schweitzer *et al.* (2010)

*Enoploclytia tepeyacensis* n. sp.

Figure 2

### 3.1. Diagnosis

Carapace elongate and cylindrical; deep cervical groove, gently inclined, intercepting dorsal margin at an angle of *ca.* 60°; dorsal midline of the carapace without intercalated plate; deep, very reduced gastro-orbital groove; deep, sinuous postcervical groove, intercepting dorsal margin at an angle of *ca.* 50°; deep, very short branchiocardiac groove, not joining postcervical groove, not extending ventrally below mid-height, intercepting dorsal margin at an angle of *ca.* 60°; deep hepatic groove, strongly convex ventrally at intersection with postcervical groove and strongly concave dorsally at intersection with cervical groove; deep inferior groove strongly convex posteriorly; all regions of the carapace covered dorsally with strong tubercles, becoming reduced in size ventrally.

### 3.2. Etymology

The trivial name alludes to El Tepeyac area, where the Mexican specimens were discovered.

Holotype. MUZ-824.

Paratype. MUZ-825.

Type locality. Jiménez (Coahuila, NE Mexico).

Geological age. Campanian (Late Cretaceous).

### 3.3. Material and measurements

Two specimens from NE Mexico (Holotype MUZ-824 - lcxp: 96.8 mm; wcxp: 53.2 mm; hcxp: 44.2 mm; Paratype MUZ-825 - lcxp: 65.9 mm; hcxp: 32.1 mm).

### 3.4. Description

Carapace elongate, cylindrical, slightly inflated between grooves; front not visible; straight dorsal margin, with a row of small, anteriorly directed spines; dorsal midline of carapace without intercalated plate; rimmed posterior margin, strongly convex medially and ventrally; ventral margin strongly curved; deep cervical groove, gently inclined, intercepting dorsal margin at an angle of *ca.* 60°; deep, subtle gastro-orbital groove, intercepting cervical groove medially; weak, short, shallow antennal groove approaches but does not intercept anterior margin; deep, sinuous postcervical groove, intercepting dorsal margin at an angle of *ca.* 50°, strongly inclined, concave forward medially, intercepting inferior groove ventrally; deep, very short branchiocardiac groove, not joining postcervical groove, not extending ventrally below mid-height, intercepting dorsal margin at an angle of *ca.* 60°; deep hepatic groove, strongly convex ventrally in intersection with postcervical groove and strongly concave dorsally in intersection with cervical groove; slightly inflated adductor muscle insertion area; deep inferior groove strongly convex posteriorly; all regions of carapace covered dorsally with strong tubercles, reducing in size ventrally; some larger tubercles arranged randomly in antennal region.

## 4. Discussion and Conclusions

As reported by Förster (1966) and Schweitzer and Feldmann (2001), representatives of *Enoploclytia* McCoy, 1849 are usually confused with those of *Palaeastacus* Bell, 1850, and *Eryma* Von Meyer, 1840. However, there are some morphological distinctions among these genera in order to justify their systematic validity within the family, as pointed out by Förster (1966). Indeed, the postcervical and branchiocardiac grooves are closely spaced and do not extend to the dorsum in *Eryma*; postcervical and branchiocardiac grooves are closely spaced and the branchiocardiac groove does not extend ventrally below mid-height of the carapace in *Enoploclytia*; postcervical and branchiocardiac grooves are longer and typically extend below the midline of the carapace in *Palaeastacus*.

According to features described by Förster (1966), the Mexican specimens are assigned to *Enoploclytia* because the short branchiocardiac groove does not extend ventrally below mid-height of the carapace. As reported by Schweitzer *et al.* (2010), five species from the Early to Late Cretaceous are known to date from the United States: *E. sculpta* Rathbun, 1926, *E. triglypta* Stenzel, 1945, *E. tumimanus* Rathbun, 1935, *E. wenoensis* Rathbun, 1935, and *E. wintoni* Stenzel, 1945. *Enoploclytia sculpta* from the Maastrichtian (Late Cretaceous) of the Ripley Fm. (Tennessee, USA), known by only pleonal somites, probably does not belong to *Enoploclytia* because it lacks



Figure 2. *Enoploclytia tepeyacensis* n. sp., early Campanian, upper Dessau Formation, Coahuila, Mexico (MUZ 825, 824). 1) Dorsal view of holotype MUZ-824; 2) Right lateral view of holotype MUZ-824; 3) Left lateral view of holotype MUZ-824; Left lateral view of paratype MUZ-825. Scale bar = 10 mm.

strong spines on the pleon (Stenzel, 1945). *Enoploclytia triglypta* from the Late Cretaceous (Coniacian) of Fannin County (Texas, USA) should be assigned to *Palaeastacus*, according to Förster (1966), because the postcervical and branchiocardiac grooves are long and typically extend below the midline of the carapace. The holotype of *Enoploclytia tumimanus* (USNM 73799) from the Paleocene of the Sucarnnoochee Shale, Midway Group (Alabama, USA) (not the Selma Chalk, Late Cretaceous, according to Rathbun, 1935) is an incomplete carapace, partially preserving the grooves, which do not allow comparison with the Mexican specimens. *Enoploclytia wenoensis* from the Early Cretaceous (Albian) of Cooke County (Texas, USA), exhibit smooth, spineless abdominal somites, and probably does not belong to *Enoploclytia*, according to Stenzel (1945). *Enoploclytia wintoni* from the Early Cretaceous (Albian) of Tarrant County (Texas, USA) is known only from right and left chelipeds, which do not allow comparison with the Mexican specimens.

Based upon the above-mentioned observations, we can attest that *Enoploclytia tepeyacensis* n. sp. is the most complete record of *Enoploclytia* known to date in North and Central America. *Enoploclytia tumimanus* Rathbun (1935) was reported for the Maastrichtian Escondido Formation of Coahuila, Mexico. Based on a pair of long chelipeds, Vega *et al.* (2007) reported *Enoploclytia gardnerae* (Rathbun, 1935) from the Paleocene Rancho Nuevo Formation, Coahuila, Mexico. They also reported *Enoploclytia* sp. for a single, poorly preserved carapace from the same unit.

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